

## 23.4 Classification of Gradually Varied Flow Profiles

It is important to systematically classify the water surface profiles in a channel before computation of flow profiles is carried out. Such classification helps to get an overall understanding of how the flow depth varies in a channel. It also helps to detect any mistakes made in the flow computation.

It may be recalled here that

$$F^2 = \frac{\alpha Q^2 T}{gA^3} \quad (23.9)$$

where  $F$  = Froude number. Substitution of Equations (23.8) and (23.9) in Equation (23.7) leads to

$$\frac{dy}{dx} = \frac{S_0 - \frac{n^2 Q^2}{A^2 R^{4/3}}}{1 - F^2} \quad (23.10)$$

For a specified value of  $Q$ , both  $F$  and  $S_f$  are functions of the depth,  $y$ . In fact, both  $F$  and  $S_f$  will decrease as  $y$  increases. Recalling the definitions for the normal depth,  $y_n$ , and the critical depth,  $y_c$ , the following inequalities can be stated

$$S_f > S_0 \quad \text{when} \quad y < y_n \quad (23.11)$$

$$S_f < S_0 \quad \text{when} \quad y > y_n$$

$$F > 1 \quad \text{when} \quad y < y_c \quad (23.12)$$

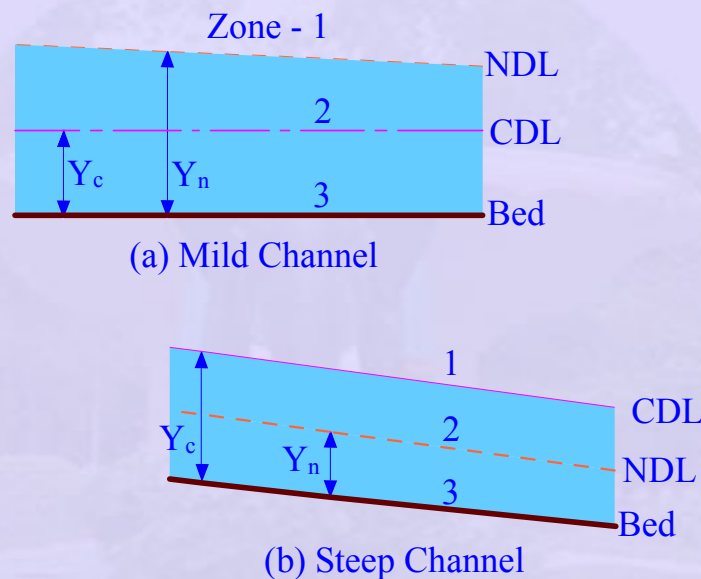
$$F < 1 \quad \text{when} \quad y > y_c$$

A gradually varied flow profile is classified based on the channel slope, and the magnitude of flow depth,  $y$  in relation to  $y_n$  and  $y_c$ . The channel slope is classified based on the relative magnitudes of the normal depth,  $y_n$  and the critical depth,  $y_c$ .

- $y_n > y_c$  : "Mild slope" (M)
- $y_n < y_c$  : "Steep slope" (S)
- $y_n = y_c$  : "Critical slope" (C)
- $S_0 = 0$  : "Horizontal slope" (H)
- $S_0 < 0$  : "Adverse slope" (A)

It may be noted here that slope is termed as "sustainable" slope when  $S_0 > 0$  because flow under uniform conditions can occur for such a channel. Slope is termed as "unsustainable" when  $S_0 \leq 0$  since uniform flow conditions can never occur in such a channel. Flow profiles associated with mild, steep, critical, horizontal, and adverse slopes are designated as M, S, C, H and A profiles, respectively.

The space above the channel bed can be divided into three zones depending upon the inequality defined by equations (23.11) and (23.12). Figure 23.2 shows these zones for a mild and a steep channel.



NDL: Normal depth line  
CDL: Critical depth line

Figure 23.2: Profile Classification

The space above both the CDL and the NDL is designated as zone-1. The space between the CDL and the NDL is designated as zone-2. The space between the channel bed and CDL/NDL (whichever is lower) is designated as zone-3. Flow profiles are finally classified based on (i) the channel slope and (ii) the zone in which they occur. For example, if the water surface lies in zone-1 in a channel with mild slope (Figure 23.3), it is designated as M1 profile. Here, M stands for a mild channel and 1 stands for zone-1.

It may be noted that an M1 profile indicates a subcritical flow since flow depth,  $y$  is greater than the critical depth,  $y_c$ .

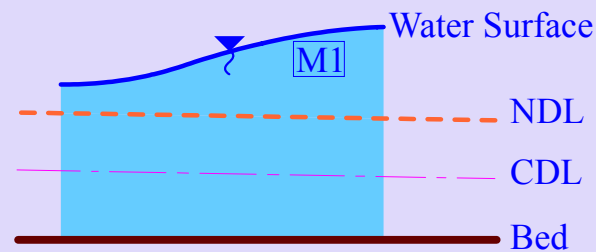


Figure 23.3: M1 Profile

Similarly, an S2 profile (Figure 23.4) indicates the water surface lies in zone-2 in a steep channel. It may be noted that a S2 profile indicates a supercritical flow since flow depth,  $y$  is lower than  $y_c$ .

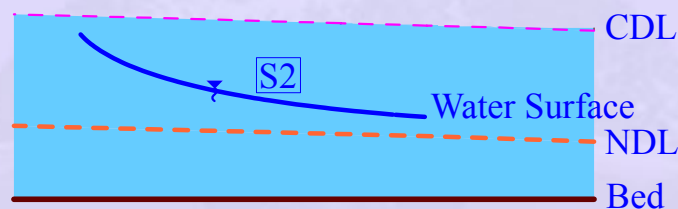


Figure 23.4: S2 Profile

Table 23.1 presents types of flow profiles in prismatic channels. In this table, a channel slope is described as critical slope when critical conditions occur for uniform flow i.e. when  $y_n = y_c$ .

Table 23.1: Types of Flow Profiles ( $S_c$ : Critical Slope)

Slope	Profile Designation			Relative position of $y$	Type of Flow
	zone - 1	zone - 2	zone - 3		
Adverse $S_0 = 0$	None	A2		$y > y_c$	Subcritical
			A3		$y < y_c$
Horizontal $S_0 = 0$	None	H2		$y > y_c$	Subcritical
			H3		$y < y_c$
Mild $0 < S_0 < S_c = 0$	M1			$y > y_n > y_c$	Subcritical
		M2		$y_n > y > y_c$	Subcritical
			M3	$y_n > y_c > y$	Supercritical
Critical $S_0 = S_c > 0$	C1	C2		$y > y_c = y_n$	Subcritical uniform -

				$y = y_c = y_n$	critical
			$C_3$	$y_c = y_n > y$	Supercritical
Steep $S_0 > S_c > 0$	S1	S2		$y > y_c > y_n$	Subcritical
				$y_c > y > y_n$	Supercritical
			$S_3$	$y_c > y_n > y$	Supercritical

## 23.5 Variation of Flow Depth

Qualitative observations about various types of water surface profiles can be made and the profile can be sketched without performing any computations. This is achieved by considering the signs of the numerator and the denominator in Equation (23.10). The following analysis helps to know (i) whether the depth increases or decreases with distance; and (ii) how the profile approaches the upstream and downstream limits. First, consider the following general points:

- $y > y_c$ ; flow is subcritical;  $F < 1$ ; denominator is positive.
- $y < y_c$ ; flow is supercritical;  $F > 1$ ; denominator is negative.
- $y = y_n$ ; flow is uniform;  $S_f = S_0$ ; numerator is zero.
- $y > y_n$ ;  $S_f < S_0$ ; numerator is positive.
- $y < y_n$ ;  $S_f > S_0$ ; numerator is negative.
- As  $y \rightarrow y_n$  ( $y$  tends to  $y_n$ );  $S_f \rightarrow S_0$ ;  $S_f \rightarrow S_0$ ; numerator approaches zero;  $\frac{dy}{dx} \rightarrow 0$ ; the surface profile approaches normal depth asymptotically.
- As  $y \rightarrow y_c$ ; Flow tends to critical conditions;  $F \rightarrow 1$ ; denominator tends to zero;  $\frac{dy}{dx} \rightarrow \infty$ ; water surface profile approaches the critical depth vertically.

It is not possible to have a vertical water-surface profile. Therefore, it is assumed that the water surface profile approaches the CDL at a very steep slope. It may be noted that when the water surface slope is very steep, it cannot be assumed that accelerations in the vertical direction are negligible. This means that the theory of gradually varied flow should breakdown in such a situation because pressure is no

longer hydrostatic in those regions. Thus equation (23.10) is not valid whenever flow depth is close to the critical depth.

As  $y \rightarrow \infty$ ;  $S_f \rightarrow 0$ ;  $F \rightarrow 0$ ;  $\frac{dy}{dx} \rightarrow S_0$ ; Water surface profile becomes horizontal as flow depth becomes very large.

For a wide channel, hydraulic mean radius  $R \approx h$  and  $F^2 = \frac{q^2}{gy^3}$ . Equation (23.10) can be simplified to

$$\frac{dy}{dx} = \frac{gy^3(S_0 y^{10/3} - q^2 n^2)}{y^{10/3}(gy^3 - q^2)}$$

where  $q$  = flow rate per unit width. It can be seen from the above equation that  $\frac{dy}{dx} \rightarrow \infty$  as  $y \rightarrow 0$ . In other words, water surface profile tends to become vertical as the flow depth tends to zero.

The qualitative characteristic of any type of water-surface profile may be studied using the points discussed earlier. For example, consider an M1 profile. For an M1 profile,

$y > y_n > y_c$ .  $y > y_c$  implies that  $F < 1$  and  $y > y_n$  implies that  $S_f < S_0$ .

Therefore,

$$\frac{dy}{dx} = \frac{S_0 - S_f}{1 - F^2} = \frac{+}{+} = +$$

This means that flow depth increases with distance  $x$ . On the downstream side, as  $y$  keeps increasing  $\frac{dy}{dx}$  tends to  $S_0$  and the water surface becomes horizontal. On the

upstream side, as  $y$  approaches the normal depth,  $y_n$ , it approaches asymptotically. The sketch of an M1 profile is shown in Figure 23.5.

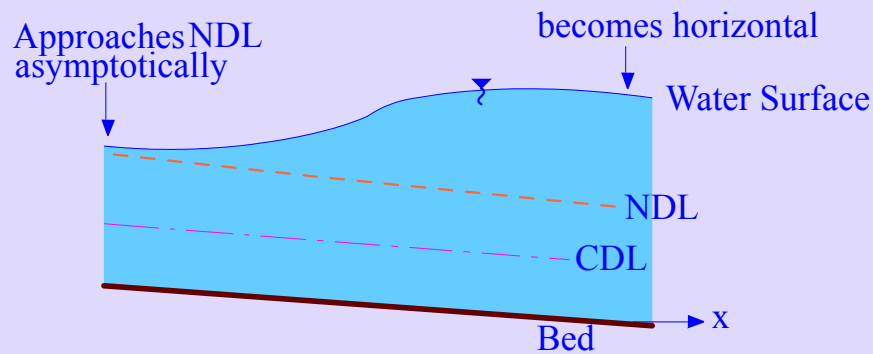


Figure 23.5: Sketch of an M1 profile

Similarly, consider an M2 profile. In an M2 profile,  $y_n > y > y_c$ .  $y > y_c$  implies that  $F < 1$  and the denominator is positive. On the other hand,  $y < y_n$  implies that  $S_f > S_0$ . Therefore,

$$\frac{dy}{dx} = \frac{S_0 - S_f}{1 - F^2} = \frac{-Ve}{+Ve} = -Ve$$

This means that flow depth decreases with distance  $x$ . On the downstream side, as the flow depth decreases and approaches the CDL, it approaches vertically. On the upstream side as the depth increases and approaches the normal depth, it approaches asymptotically. The sketch of an M2 profile is shown in Figure 23.6.

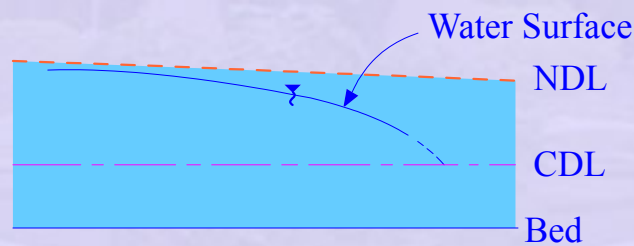


Figure 23.6: Sketch of an M2 profile

Now, Consider an S2 profile. In an S2 profile,  $y_c > y > y_n$ .  $y < y_c$  implies that  $F > 1$  and the denominator is negative.  $y > y_n$  implies that  $S_f < S_0$ . Therefore,

$$\frac{dy}{dx} = \frac{S_0 - S_f}{1 - F^2} = \frac{+Ve}{-Ve} = -Ve$$

This means that flow depth decreases with distance  $x$ . On the downstream side, as  $y$  decreases towards  $y_n$  it approaches NDL asymptotically. On the upstream side, as  $y$  increases toward  $y_c$ , it approaches CDL almost vertically. The sketch of an S2 profile is shown in Figure 23.7.

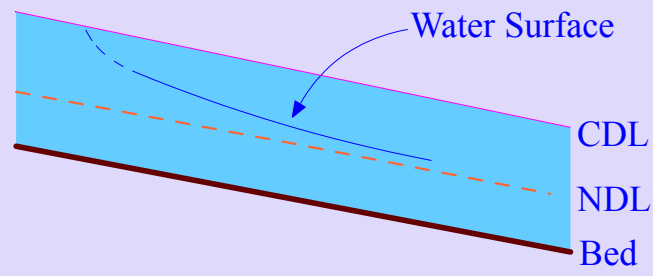
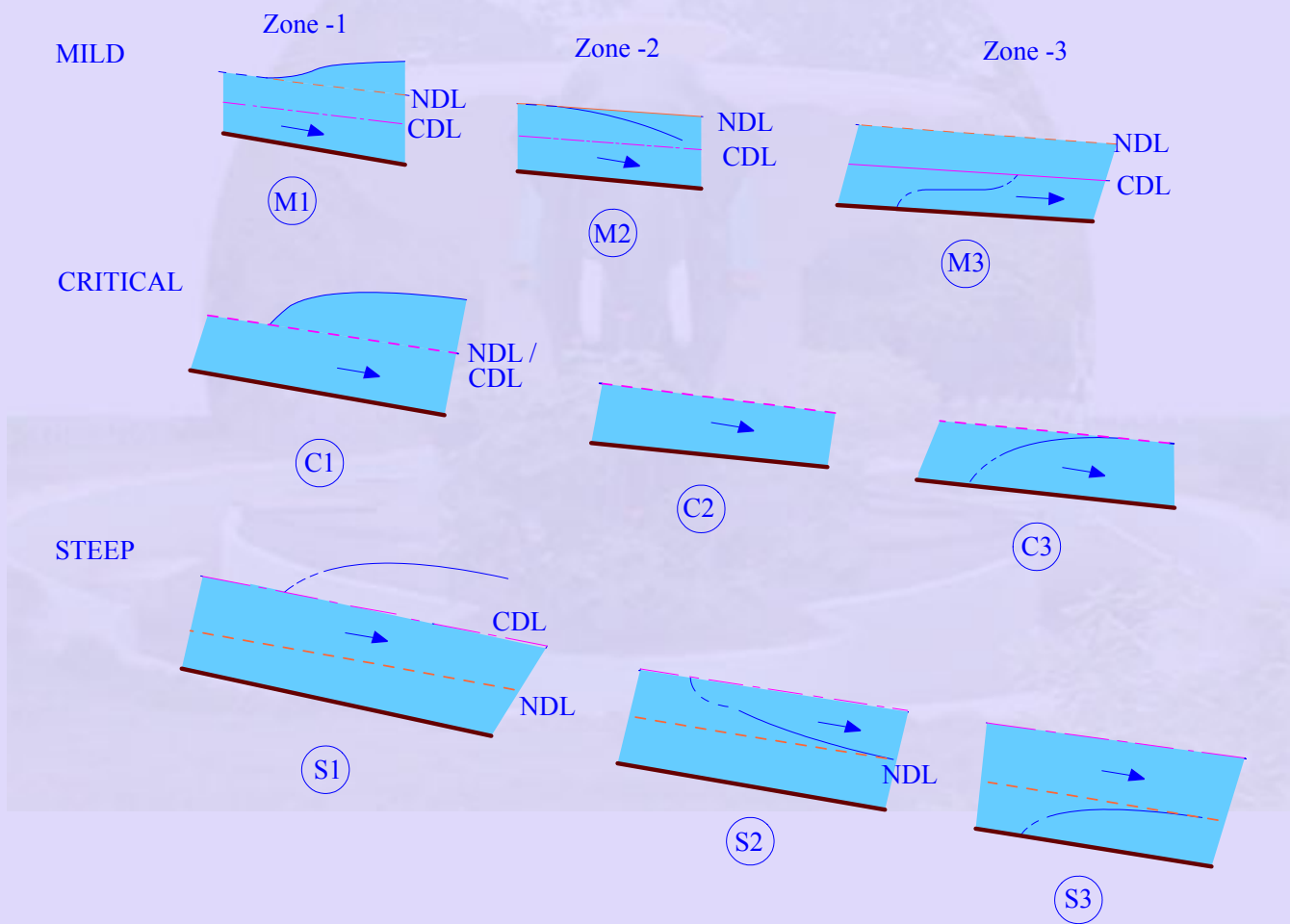


Figure 23.7: Sketch of an S2 profile

Proceeding in a similar manner, other water surface profiles can be sketched. These sketches are shown in Figure 23.8. The profiles are shown in dashed lines as they approach the CDL and the channel bed to indicate that gradually varied flow assumption is not valid in those regions.



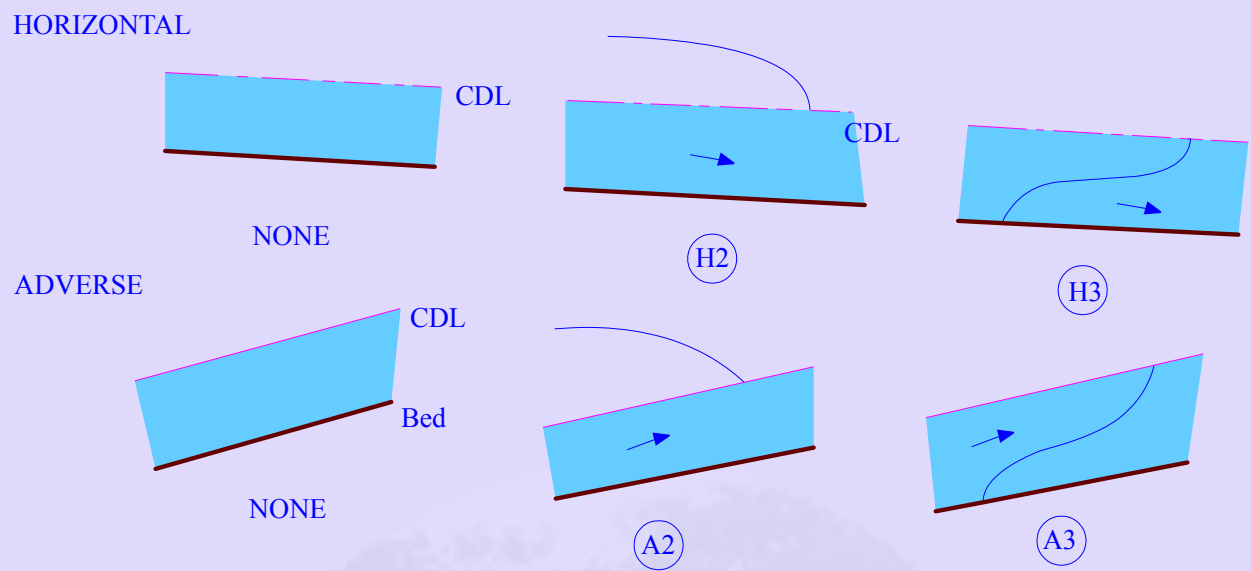


Figure 23.8: Water Surface Profiles

